MODELING WATER MOVEMENT THROUGH A FLORIDA SPODOSOL AND ENTISOL USING HYDRUS-2D Davie Kadyampakeni¹, Kelly Morgan¹, Peter Nkedi-Kizza² and Arnold Schumann³ UNIVERSITY of **FLORIDA**

¹University of Florida, Soil and Water Science Department, Southwest Florida Research and Education Center, 2686 SR 29N, Immokalee, FL 34142, USA. ²University of Florida, Soil and Water Science Department, 2169 McCarty Hall, P.O. Box 110290, Gainesville, FL 32611, USA. ³University of Florida, Soil and Water Science Department, Citrus Research and Education Center, 700 Experiment Station Road, Lake Alfred, FL 33850, USA.



To measure and predict water movement patterns within a 0.5-m radius by 0.6-m deep simulation domain using drip- and microsprinkler irrigation. Governing Equations The predictive equation for the unsaturated hydraul function in terms of soil water retention parameters given by van Genuchten (1980) as: $\theta(h) = \frac{\theta_r + \frac{\theta_s - \theta_r}{[1 + \alpha h ^n]^m}}{\theta_s} h < 0$		Objective
Governing EquationsThe predictive equation for the unsaturated hydraulfunction in terms of soil water retention parametersgiven by van Genuchten (1980) as: $\theta(h) = \frac{\theta_r + \frac{\theta_s - \theta_r}{[1 + \alpha h ^n]^m}}{\theta_s} h < 0$ θ_s	T p s n	o measure and predict water movement atterns within a 0.5-m radius by 0.6-m deep imulation domain using drip- and nicrosprinkler irrigation.
The predictive equation for the unsaturated hydraul function in terms of soil water retention parameters given by van Genuchten (1980) as: $\theta(h) = \frac{\theta_r + \frac{\theta_s - \theta_r}{[1 + \alpha h ^n]^m}}{\theta_s} h < 0$		Governing Equations
	ן נ	The predictive equation for the unsaturated hydraul function in terms of soil water retention parameters given by van Genuchten (1980) as: $\theta(h) = \frac{\theta_r + \frac{\theta_s - \theta_r}{[1 + \alpha h ^n]^m}}{\theta_s} h < 0$

2) near the Citrus Research and Education Center, Lake Alfred, Fla. (28°5' N, 81°45' W) on the Florida Ridge.

Irrigation treatments

Treatments were as follows:						
(1) Conventional microsprinkler practice (CMP) –						
irrigated weekly;						
(2) Drip open hydroponics system (DOHS) –						
irrigated daily in small pulses;						
(3) Microsprinkler open hydroponics system						
(MOHS) – irrigated daily.						

Estimation of soil moisture

•Determination of soil water release curves (SWRC) at 0 through 100 kPa.

Candler	OBS vs. MS –spring at 40 cm	0.87
Candler	OBS vs. DRIP-spring at 10 cm	0.99
Candler	OBS vs. DRIP-spring at 40 cm	0.93
Candler	DRIP vs. MS at 10 cm	1.00
Candler	DRIP vs. MS at 40 cm	1.00
Immokalee	OBS vs. MS-spring at 10 cm	0.99
Immokalee	OBS vs. DRIP-spring at 10 cm	1.00
Immokalee	OBS vs. MS-spring at 40cm	1.00
Immokalee	OBS vs. DRIP-spring at 40 cm	0.95
Immokalee	DRIP vs. MS-spring at 10 cm	1.00
Immokalee	Drip vs. MS-spring at 40 cm	0.99
Immokalee	OBS vs. MS-summer at 10 cm	0.99
Immokalee	OBS vs. DRIP-summer at 10 cm	0.96
Immokalee	OBS vs. MS-summer at 40cm	0.99
Immokalee	OBS vs. DRIP-summer at 40 cm	1.00

[¶]OBS-Observed or measured in the field, MS-Microsprinkler irrigation, DRIP-Drip irrigation, R²-Coefficient of determination

Where $m = 1 - \frac{1}{n}, n > 1$ $S_e = \frac{(\theta - \theta r)}{(\theta_s - \theta_r)}$ Where θ_r , θ_s , K_s and *l* are residual water content (L³L⁻ ³), saturated water content (L³L⁻³), saturated hydraulic conductivity (LT⁻¹), and pore connectivity. α (L⁻¹) and *n* are empirical coefficients affecting the shape of the hydraulic functions. The governing flow equations for water flow and nutrient transport are given by the Richards (1931) $\frac{\partial \theta}{\partial t} = \frac{\partial}{\partial x_i} \left[K \left(K_{ij}^A \frac{\partial h}{\partial x_i} + K_{iz}^A \right) \right] - s(h)$ Where θ is the volumetric water content [L³L⁻³], h is the pressure head [L], , x_i (i=1, 2) are the spatial

•Gravimetric and sensor-based measurement of soil moisture content.

•Use of Br tracer for monitoring water movement.

•Calibration of HYDRUS-2D using site-specific data.

•Simulation of water movement through a 0.5-m radius by 0.6-m deep simulation domain.

Discussion and Conclusions

•Results indicate reasonably good agreements between measured and predicted values water content ($\mathbb{R}^2 > 0.87$).

Table 2: Statistical comparison between observed and simulated Br contents in spring and summer 2011.

			RMSE (mg kg ⁻
Soil	¶Comparison	R ²	1)
Candler	OBS vs. MS –spring at 10 cm	0.89	0.18
Candler	OBS vs. MS –spring at 40 cm	0.76	0.25
Candler	OBS vs. DRIP-spring at 10 cm	0.96	0.35
Candler	OBS vs. DRIP-spring at 40 cm	0.75	0.46
Immokalee	OBS vs. MS-spring at 10 cm	0.79	0.57
Immokalee	OBS vs. DRIP-spring at 10 cm	0.90	0.44
Immokalee	OBS vs. MS-spring at 40cm	0.74	0.06
Immokalee	OBS vs. DRIP-spring at 40 cm	0.63	0.04

OBS-Observed or measured in the field, MS-Microsprinkler irrigation, DRIP-Drip irrigation, R²-Coefficient of determination, **RMSE-Root mean square error**



medium is isotropic), K is the unsaturated hydraulic conductivity function (LT⁻¹), and s is a sink/source term [L³L⁻³T⁻¹], accounting for root water uptake (transpiration). The sink/source represents the volume of water removed per unit time from a unit volume of soil due to compensated citrus water uptake.

coordinates [L] for two-dimensional flow, t is time [T],

K^A_{ii} are components of a dimensionless anisotropy

tensor K^A (which reduces to the unit matrix when the

•Br movement was also well predicted ($R^2 >$ 0.87 and RMSE 0.04 -0.46 mg/kg)

•The results suggest that a carefully calibrated **HYDRUS-2D model could be used for** irrigation decision support on Florida's **Spodosols and Entisols.**

Richards, L.A. 1931. Capillary conduction of fluid through porous medium. Physics 1:318-333.

VanGenuchten, M. 1980. A closed-form equation for predicting the hydraulic conductivity of unsaturated soils. Soil Sci. Soc. Am. J. 44:892-898.