

An intercomparison study of TSM, SEBS and SEBAL using high resolution imagery and Lysimetric data



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INTRODUCTION

Evapotranspiration (ET) mapping has many applications including crop water management, climate change impact assessment, hydrological modeling, groundwater recharge studies, irrigation performance, and land use planning. Satellite-based thermal infrared remote sensing has greatly contributed to the development and improvement of remote sensing-based evapotranspiration (RS-ET) mapping algorithms. Testing and validation of RS-ET algorithms across a range of hydrometeorological and surface cover conditions is important to fill in the existing gap in the operationalization of these algorithms. The primary objective of this research was to test and improve three widely used RS-ET models. Three algorithms evaluated in this study include: SEBAL (Surface Energy Balance Algorithm for Land), SEBS (Surface Energy Balance System) and TSM (Two Source Energy Balance Model).

RESULTS

Table 1: Performance statistics for T_s (Obs. Mean: 34.6°C), R_n (Obs. Mean: 577), and G_o (Obs. Mean:37); H (Obs. Mean:182); (no. of observations = 40).

Estimated parameter	Mean	MBE	%MBE	RMSE	%RMSE	MAE	MAPD	NSE	R ²
T _s (°C)	34.8	0.2	0.7	1.6	4.5	1.1	3.2	0.94	0.9
$R_n (W m^{-2})$	571	-5.6	-1.0	29.5	5.1	23.5	4.1	0.72	0.7
G _o (W m ⁻²)	35	-2.0	-5.3	16.8	45.2	13.2	35.4	0.21	0.2
H _{SEBAL}	228	57.6	33.9	88.7	52.2	72.2	42.5	0.26	0.6
H _{SEBS}	190	8.2	4.5	53.9	29.6	45.2	24.8	0.78	0.8
H _{TSM}	121	-60.7	-33.4	100.5	55.2	83.5	45.8	0.24	0.5

THEORY

SEBAL, SEBS, and TSM utilize the Monin-Obukhov similarity (MOS) theory to solve for the sensible heat and calculate ET as the residual of the surface energy balance (Net Radiation = Soil Heat Flux + Sensible Heat Flux + Latent Heat Flux). In general, the residual surface energy balance scheme can be categorized into singlesource (SEBS and SEBAL) and dual-source (TSM) model, differing in their treatment of soil and vegetation source contribution as composite or distributed, respectively. Below is the bulk formulation of sensible heat (H) based on the gradient-resistance relation as defined by each of the three algorithms.





 $ET_{estimated} = 1.0 ET_{observed} - 0.08$

 $R^2 = 0.74$

1.00

0.80

E 0.60



Table 2: Performance statistics for instantaneous ET (mm h⁻¹) for the complete data set, irrigated and dryland fields. Observed mean for the complete dataset, irrigated and dryland fields were 0.54, 0.66 and 0.42 mm h⁻¹ respectively.

Models	n	Mean	MBE	%MBE	RMSE	%RMSE	MAE	MAPD	NSE
SEBAL	40^{\bullet}	0.48	-0.08	-14.1	0.14	25.9	0.12	21.5	0.50
	20▲	0.63	-0.02	-3.1	0.09	13.4	0.08	11.5	0.81
	20 •	0.30	-0.14	-31.8	0.18	41.3	0.17	37.8	-0.55
SEBS	40^{\bullet}	0.53	-0.01	-1.0	0.08	15.7	0.07	12.5	0.85
	20▲	0.67	0.01	2.1	0.09	14.1	0.08	11.5	0.79
	20 •	0.40	-0.02	-5.9	0.08	18.1	0.06	14.1	0.78
TSM	40 [•]	0.62	0.08	15.1	0.12	22.2	0.09	17.4	0.70
	20▲	0.69	0.03	5.1	0.07	11.3	0.05	8.5	0.86
	20 [•]	0.55	0.13	31.4	0.15	36.0	0.13	31.4	0.12



is specific heat of air at constant pressure $(Jkg^{-1}K^{-1})$ and $r_{ah,1,2}$ is aerodynamic resistance (sm⁻¹) between two near-surface heights, z_1 and z_2 taken as 0.1 and 2 m. The dT parameter (K) represents the near-surface temperature difference between z_1 and z_2 .

Pitfall: The parameter dT is assumed to be a linear function of radiometric surface temperature across the study region. The dT function is derived from two point known as the hot and cold pixel whose selection is highly subjective to analysts' decision.

 $H_{\rm S} = \rho_{\rm a} \, C_{\rm p} \, \frac{T_{\rm S} - T_{\rm AC}}{T_{\rm s}}$ where θ_0 is the potential surface temperature and θ_a is $H = \rho_a C_p \frac{T_{AC} - T_A}{r}$ the potential air temperature. The aerodynamic resistance r_{ah} , is defined as the resistance from height $z_{oh}+d_o$ (d_o is zero where H_C , H_S and H are the sensible heat for canopy, soil and total. T_C , T_A , plane displacement height, and $T_{\rm S}$, and $T_{\rm AC}$ are the temperatures of the z_{oh} (m) is roughness length for canopy, air, soil and air temperature heat transport) having an within the canopy boundary layer. r_x is aerodynamic temperature, to the resistance in the boundary layer near the height z_{ref}. z_{oh} is related to the canopy, r_s is the resistance to heat an excess resistance parameter flux in the boundary layer immediately above the soil surface and r_a is the **Pitfall:** Highly sensitive to the aerodynamic resistance. air temperature input data. **Pitfall**: Relatively highly parameterized hence requires more and accurate crop specific information.

Observed ET (mm h⁻¹)

 $ET_{estimated} = 0.70 ET_{observed} + 0.24$ $R^2 = 0.88$ **(C)**

1.00

0.80

[•]Complete data, [•]Irrigated fields data, [•]Dryland fields data

SUMMARY

1. Performance statistics for T_s , R_n , and G_o for the complete data set showed good agreement against the measured data.

2. SEBS performance was superior, as evidenced by the smaller error indices and absence of bias error in both dryland and irrigated fields.

3. SEBAL under estimated ET with large variance in the individual errors and poor performance for dryland conditions. TSM over estimated dryland ET and had significant bias error, however, overall performance of TSM was better than SEBAL.

Results suggest that all three models have the potential to be developed as an operational tool for managing water resources by providing accurate and economical spatial ET information.

MATERIALS AND METHODS

 kB^{-1} .

The algorithms were executed on 10 high resolution airborne images Bushland acquired during the and Agricultural Evapotranspiration Remote Sensing Experiment 2007 2008 (BEAREX07 and and BEAREX08) field campaign and ET validated against hourly measurements four from large precision weighing lysimeters, two each placed in irrigated and dryland fields. Images were acquired for tall (Corn and Sorghum) and short (Cotton) crops from early to midcropping season representing diverse set of agricultural surface roughness with varied land surface energy balance systems.



1.00

0.80

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